

Holocene anthropogenic landscapes in the Balkans: the palaeobotanical evidence from southwestern Bulgaria

Elena Marinova · Spassimir Tonkov ·
Elissaveta Bozilova · Ivan Vajsov

Received: 11 April 2011 / Accepted: 23 December 2011
© Springer-Verlag 2011

Abstract Palaeoecological reconstructions from the region of southwestern Bulgaria were used for inferring the human impact on the vegetation and landscape during the last 8 millennia. They are based on data from pollen analyses of lakes and peat-bogs, plant macrofossils, archaeobotanical finds and radiocarbon dating. During the early Holocene, after 7900 cal. B.P. (5950 cal. B.C.) the climate changed to cooler summers, milder winters and higher precipitation resulting in the formation of a coniferous belt dominated by *Pinus* sp. and *Abies alba*. These favorable environmental pre-conditions had a positive influence on the Neolithisation of the Balkans after the 8200 cal. B.P. (6250 cal. B.C.) cold event, which caused drought in the Eastern Mediterranean. Direct evidence from wood charcoal records from the Neolithic settlement layers in the study area shows a slight modification of the surrounding woodlands and an increase of the light-demanding components, probably

expressed through larger forest border zones and thinning out of the wood stands. The increase in the number of settlements in the valleys of southwestern Bulgaria intensified the human activity visible in the palaeobotanical record from 6950 cal. B.P. (5000 cal. B.C.) onwards. Between ca. 5700–5100 cal. B.P. (3800–3200 cal. B.C.) signs of anthropogenic influence on the vegetation are virtually absent. The intensity of human impact increased notably after 3200 cal. B.P. (1400–1250 cal. B.C., approx. Late Bronze Age), documented by a rise of pollen anthropogenic indicators. The final transformations in the natural forest cover after 2750 cal. B.P. (800 cal. B.C. onset of the Iron Age) marked the reduction of the coniferous forests dominated by *Abies alba* and *Pinus* sp. and the expansion of *Fagus sylvatica* and *Picea abies*. These vegetation changes are contemporaneous with increase of the palaeofire activities and the next peak of anthropogenic indicators. The changes in the landscape during the Roman period and the medieval period reflect regional environmental features and were forced by the diversification of anthropogenic activity.

Keywords Pollen · Wood charcoal analysis · Human impact · Climate change · South Eastern Europe

Introduction

Palaeobotanical proxy data can give a comprehensive view of anthropogenic impact on the natural vegetation, providing data on human-environment interactions. In combination with archaeological data, palaeoecological information allows the reconstruction of interactions and/or adaptation of past societies to Holocene climate changes during different historical epochs (Gaillard 2007).

Communicated by W. Kirleis.

Electronic supplementary material The online version of this article (doi:10.1007/s00334-011-0345-8) contains supplementary material, which is available to authorized users.

E. Marinova (✉)
Center for Archaeological Sciences, Katholieke Universiteit
Leuven, Celestijnenlaan 200E, 3001 Leuven, Belgium
e-mail: elena_marinova@gmx.de;
Elena.Marinova@bio.kuleuven.be

S. Tonkov · E. Bozilova
Faculty of Biology, Department of Botany, Laboratory
of Palynology, Sofia University St. Kliment Ohridski,
8 Dragan Tzankov blvd, 1164 Sofia, Bulgaria

I. Vajsov
National Archaeological Institute with Museum, Bulgarian
Academy of Sciences, Saborna 2, 1000 Sofia, Bulgaria

Southwestern Bulgaria is of special interest for studies on this topic as its natural vegetation has been influenced by anthropogenic activity since at least 8,500 years ago, starting with the introduction of Neolithic farming and continuously increasing through millennia of human occupation. The area is considered as one of the routes for the Neolithisation of southeastern Europe, playing a key role in the prehistory of the region because of its special geographic position linked with the Aegean in the south, the Thracian plain to the east and the Danube valley to the north. This position of the study area makes it very suitable for tracing out various regional land use developments and influences on the vegetation. Its natural conditions have a transitional character from Mediterranean to continental climate and vegetation, and could be crucial for understanding the linkage between the near-eastern and European land use practices.

The main periods of human occupation in the region, if considered against the background of the main palaeoclimatic events during the Holocene, show certain parallels and relationships (for discussion of some of them see Weninger et al. 2009). The Neolithisation in the study area was preceded by the 8200 cal. B.P. (ca. 6200 cal. B.C.) cooling event which influenced the environment and human occupation in the study area (Weninger et al. 2006). Another period of importance is the end of the Holocene climatic optimum, which can be roughly associated with Bond cycle 4: 5400–4800 cal. B.P. (3450–2850 cal. B.C.) (Bond et al. 2001) and which is more or less contemporary with the transitional period between the end of the Chalcolithic period in Bulgaria (ca. 3800–3500 cal. B.C.) and the onset of the Early Bronze Age (3000–2800 cal. B.C.) (Görsdorf and Boyadzhiev 1996). The climatic fluctuations that followed, considered as global, such as Bond Event 3 or the 4200 cal. B.P. (2250 cal. B.C.) event, correspond to the second half of the Early Bronze Age, which was a period of low settlement activity in the study area (Grębska-Kulowa and Kulow 2007). The evidence for the adjacent areas, such as the Aegean and Anatolia (Roberts et al. 2008, 2011), indicate aridity for this period. The rather arid period around 2850 cal. B.P. (900 cal. B.C., corresponding to Bond event 2), was contemporary with the Early Iron Age in the region. The last fairly well globally recorded palaeoclimatic event (Wanner et al. 2008) is the period of ca. 700–500 cal. B.P. (A.D. 1300–1500) corresponding to the Little Ice Age.

In order to evaluate the degree of human impact in the region under consideration, a number of investigations combining pollen analysis, archaeobotanical studies, radiocarbon dating, and information on past climate changes were carried out over a period of 20 years. These studies focused on the introduction of agriculture, the practice of stock-breeding and the collection of plant material for various purposes, as well as the human-environment

interactions (e.g. Popova and Bozhilova 1997; Marinova et al. 2002; Kreuz et al. 2005; Marinova 2006; Bozilova and Tonkov 2007; Marinova and Thiebault 2008; Marinova and Popova 2008).

Since at present this information is spread over different publications and partly unpublished, this paper provides the first comprehensive overview and critical evaluation of past vegetation changes in relation to human occupation for the region. As such, it presents evidence on the long term development of the human landscapes in southwestern Bulgaria during the Holocene. The main sources of information for this study are the palynological and anthracological data from south western Bulgaria. Together with this some relevant archaeobotanical information on seed/fruit remains of crops and weeds is also taken into consideration. Possible interpretations for natural and anthropogenic landscape changes are discussed, and the most important stages of these changes are defined. The emphasis will finally be put into a regional context, by comparison with data from the adjacent areas.

Study area

The study area is situated in southeastern Europe and can be divided in two sub-regions corresponding to the main geographical features of southwestern Bulgaria: (1) the Upper and Middle-Struma river valley and its adjoining slopes, and (2) the mountain areas; to the east of the valley the Rila and Pirin mountains, and to the west several smaller mountain ranges (Konjavaska, Osogovo and Maleshevska). The locations of the palynological and archaeological sites considered here are shown in Fig. 1.

Natural environment

The climate in the study area, especially its southern part, is under the influence of the Mediterranean, while more continental conditions prevail further north. In general, it can be considered as being transitional between continental and sub-Mediterranean climate. The mean annual temperature in the lowlands varies from 14°C in the south to 10°C in the north. The annual precipitation in the lowlands is from 780 mm in the south to ca. 600 mm in the north. Above 1,000 m the climate changes to mountainous and the mean temperature drops by 0.5°C with each 100 m increase in altitude. The mean annual precipitation is 800–1,000 mm, much of it snow at higher altitudes (Kopravev 2002).

According to Velchev (2002) the natural vegetation in the lowlands of the study area (up to 900 m) is composed of xerothermic and xeromesophilous oak and hornbeam forests dominated by *Quercus cerris*, *Q. pubescens*, *Q. frainetto*, *Q. dalechampii*, *Carpinus orientalis* and *C. betulus*, with an admixture of *Ulmus glabra*, *Acer platanoides* and

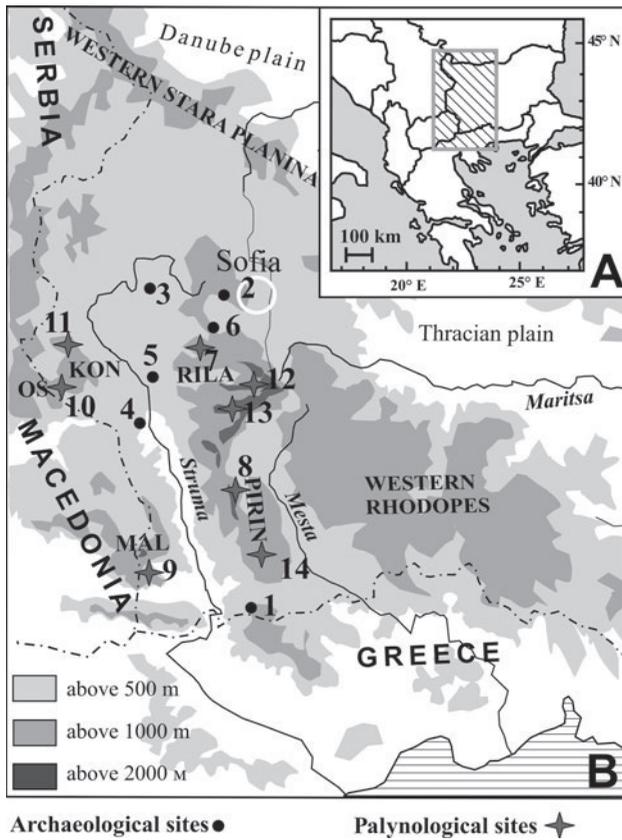


Fig. 1 a Overview map with the study area indicated by a hatched rectangle. b Detailed map showing the palynological sites under consideration, the three additionally included in the summary given in Fig. 4, and the archaeological sites; KON Konjavska Mountain, OS Osogovo Mountain, MAL Maleshevska Mountain. Archaeological sites: 1 Kovacevo 172 m a.s.l., 2 Slatina 550 m a.s.l., 3 Galabnik 650 m a.s.l., 4 Balgarchevo 368 m a.s.l., 5 Slatino 405 m a.s.l., 6 Kremenik 720 m a.s.l.; Palynological sites: 7 Lake Trilistnika 2,216 m, 8 Lake Ribno Banderishko 2,190 m a.s.l., 9 Peat bog Maleshevska 1,700 m a.s.l., 10 Peat bog Osogovo 1,720 m, 11 Tshokljovo marsh 870 m a.s.l., 12 Ostrezki Lakes 2,330 m a.s.l., 13 Dry Lake 1,900 m a.s.l., 14 Peat bog Mutorog 1,700 m a.s.l

dominated by *Pinus sylvestris*, *Picea abies* and *Pinus peuce*. The subalpine area (2,000–2,400 m) in the Rila and Pirin mountains is dominated by a thick, impenetrable formation of *Pinus mugo* with *Juniperus sibirica* and *Vaccinium myrtillus*. Above it the alpine area is occupied by various herb communities.

Archaeological background

The development of the prehistoric cultures in southwestern Bulgaria, and particularly along the Struma valley, is a key question in Balkan prehistory. The direct territorial connection of this region with the northern Aegean coast, and from there with Anatolia, influenced the specific dynamics of cultural changes through all prehistoric periods (Pernicheva 1995; Nikolov 2007). The Struma valley is considered as one of the primary routes for the Neolithisation of the Balkan Peninsula (Lichardus-Itten et al. 2006) which according to the available archaeological evidence and radiocarbon dates started around 6200/6100 cal. B.C. (Görsdorf and Bojadziev 1996; Boyadzhiev 2009). The Early Neolithic settlements were situated in the foothills of the mountains, between 400 and 650 m. Quite probably their location reflects favorable climatic and ecological conditions (Todorova and Vaisov 1989). An increase in the number of prehistoric settlements is observed during the Neolithic with a maximum around the last quarter of the 6th millennium B.C. (Grębska-Kulowa and Kulow 2007). During the second half of the Late Neolithic (5200–4900 cal. B.C.) new, larger settlements with surface areas of up to 16 ha were founded. The number of the settlements increased considerably during this period, indicating a strong increase in the population (Grębska-Kulowa 2005). By the end of the Late Neolithic and the beginning of the Early Chalcolithic (4900–4850 cal. B.C.) the number of settlements was decreasing and this also continued during the Early Bronze Age (3200–2500 cal. B.C.). The settlements were of limited size and were usually surrounded with palisades or stone walls. This resulted in a chain of protected sites in the Struma valley. A second maximum in the settlement activities in the area was observed during the Late Bronze Age (1400–1200 cal. B.C.) (Grębska-Kulowa and Kulow 2007). Between the Early Bronze Age and the beginning of the Late Bronze Age there is a chronological hiatus of about 600 years from the duration of which no particular sites and cultural content have yet been found (Boyadzhiev 2007).

Subsequently, settlements were founded in the lower parts of the Struma valley, near the river, only during the Late Iron Age probably as a result of the increase in population. During the 6th and in the first half of the 5th century B.C. the broader geographical area around the lower courses of the Mesta (Nestos), Struma (Strymon) and

159 *A. pseudoplatanus*. The sub-Mediterranean floristic elements in these forests increase from north to south. Today
160 human impact has turned many of those forests into secondary communities dominated by sub-Mediterranean
161 species like *Carpinus orientalis* and *Paliurus spina-christi*, and partly by *Juniperus oxycedrus*, *Phillyrea latifolia*,
162 and partly by *Quercus coccifera* und *Pistacia terebinthus*. The open habitats created by anthropogenic activities are usually
163 dominated by xerophylous shrubs and herbs mainly of a sub-Mediterranean or steppe character such as *Genista
164 rumelica*, *Astragalua angustifolius*, *Amygdalus nana*,
165 *Artemisia alba*, *Agropyron brandsae* etc. (Bondev 1991).

166 In the mountain ranges to the west of the Struma valley the zone between 1,000 and 1,600 m a.s.l. is dominated
167 by forests of *Fagus sylvatica* while in the Rila and Pirin mountains above 1,500 m a coniferous belt exists
168
169
170
171
172
173
174

226 Vardar (Axios) rivers experienced remarkable economical,
 227 political and cultural prosperity. During the 5th–4th centu-
 228 ry B.C. the area was integrated in the Kingdom of Mac-
 229 edonia, and later on, in the 1st century A.D. it became part
 230 of the Roman Empire (Delev 2002). Of importance for the
 231 land use strategies is also the fact that woodland-consum-
 232 ing iron production was practiced on a large scale in the
 233 mountainous area between the rivers Mesta and Struma
 234 during the Medieval. Recent surveys (Athanasov et al.
 235 2010) have found iron slag on numerous sites in the area
 236 since from even Late Antiquity. In general, the study area
 237 was also involved in transhumance at least from the Late
 238 Bronze Age (Leshtakov 2006) until the First World War.
 239 This should also be considered when interpreting the pal-
 240 aeocological evidence for the past landscapes.

241 **Materials and methods**

242 The evidence used for this study was taken first from several
 243 palynological profiles considered most appropriate due to
 244 their relatively good chronological framework and location.
 245 These data [summarized in Fig. 2 and ESM (Electronic
 246 Supplementary Material) Figs 1–5] were combined with
 247 anthracological evidence gathered from archaeological sites
 248 (Fig. 3). The combination and comparison of data coming
 249 from natural sediments (pollen) and anthropogenic layers
 250 (wood charcoals) allow the creation of a more detailed and

complete picture of the past anthropogenic changes in the
 vegetation. Moreover, in interpreting the palynological and
 anthracological data, the archaeobotanical evidence on the
 ancient agriculture and land use in the study area was also
 taken into consideration (see overview in Valamoti 2004;
 Borojevic 2006; Marinova 2006; Popova 2009).

Finally, an attempt was made to correlate the palaeobo-
 tanical evidence from the study area with information on
 global rapid climate changes as given by Mayewski et al.
 (2004), and with the available radiocarbon datasets
 and archaeological chronology provided by Chohadzhiev
 (2007), Görsdorf and Bojazhiev (1996), Vajsov (1998),
 Gatsov and Boyadziev (2009) and Stefanovich and Bankoff
 (1998). The correlation is presented in Fig. 4.

Palynological data

The palynological data used in this paper originate from
 investigations already published in Tonkov (1988, 2003),
 Tonkov et al. (2002, 2008) and Tonkov and Bozilova
 (1992a, b). Simplified percentage pollen diagrams with
 selected taxa of importance for the current study are given
 as ESM Figs 1–5. The exaggeration of the curves is marked
 by a dotted pattern and is 10× in all of the diagrams. In all
 these detailed pollen diagrams the sum of the arboreal
 pollen averages 450–500 grains, with the exception of
 Tshokljovo Marsh where it was ca. 300 arboreal grains. For
 the majority of the pollen types the nomenclature of Beug

Fig. 2 Correlation between the local pollen assemblage zones in the pollen diagrams of southwestern Bulgaria (KON Konjavaska Mountain, OS Osogovo Mountain, MAL Maleshevska Mountain)

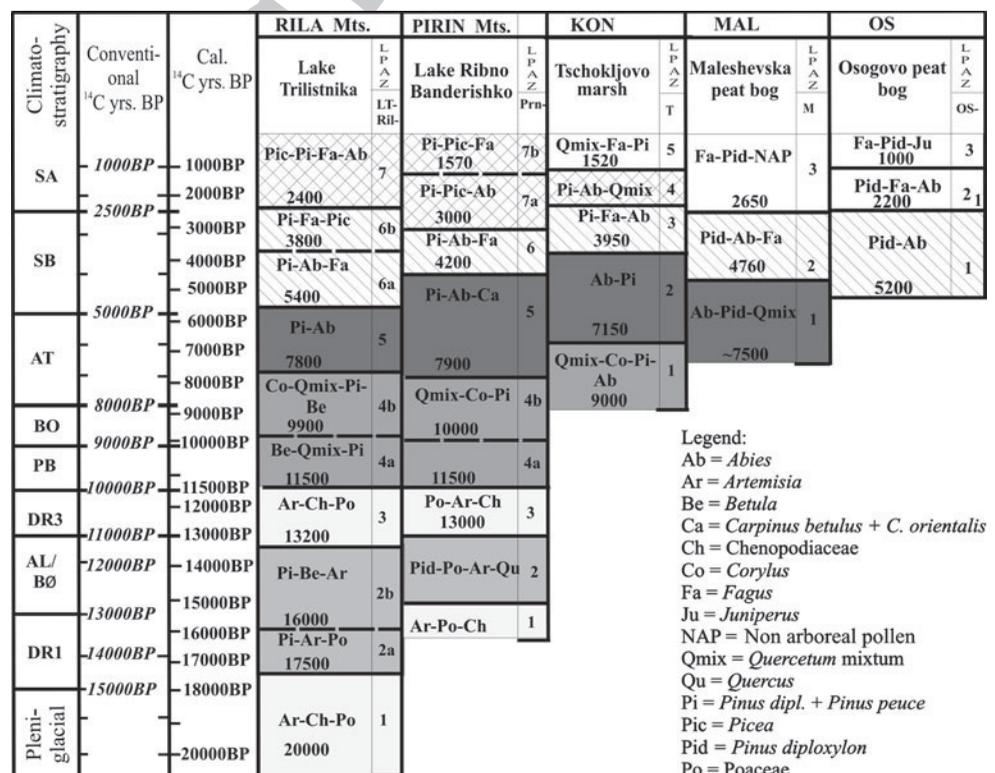


Fig. 3 Wood charcoal records from archaeological sites in the study area presented in percentage values of the numbers of studied wood charcoal fragments

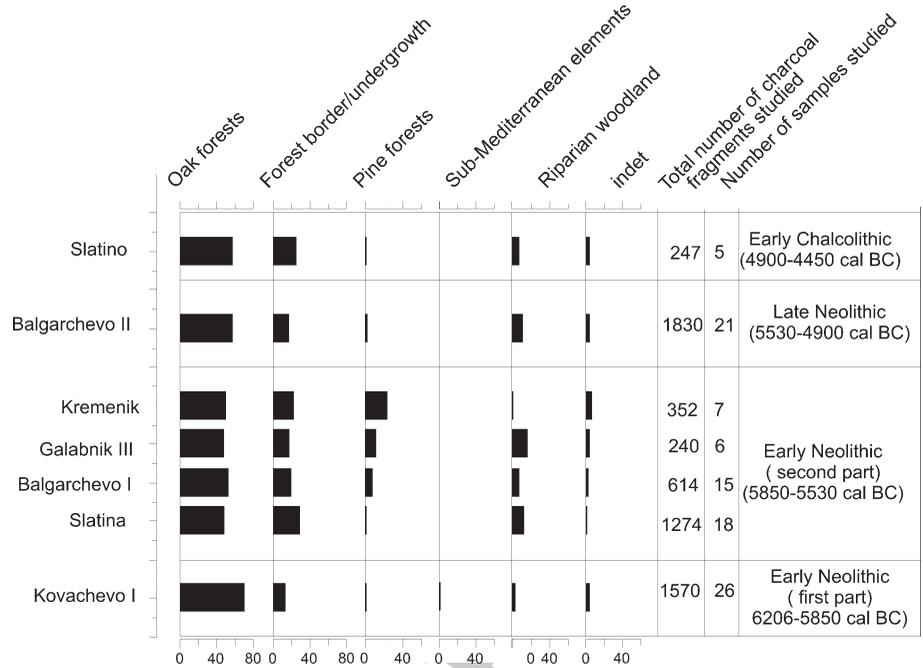
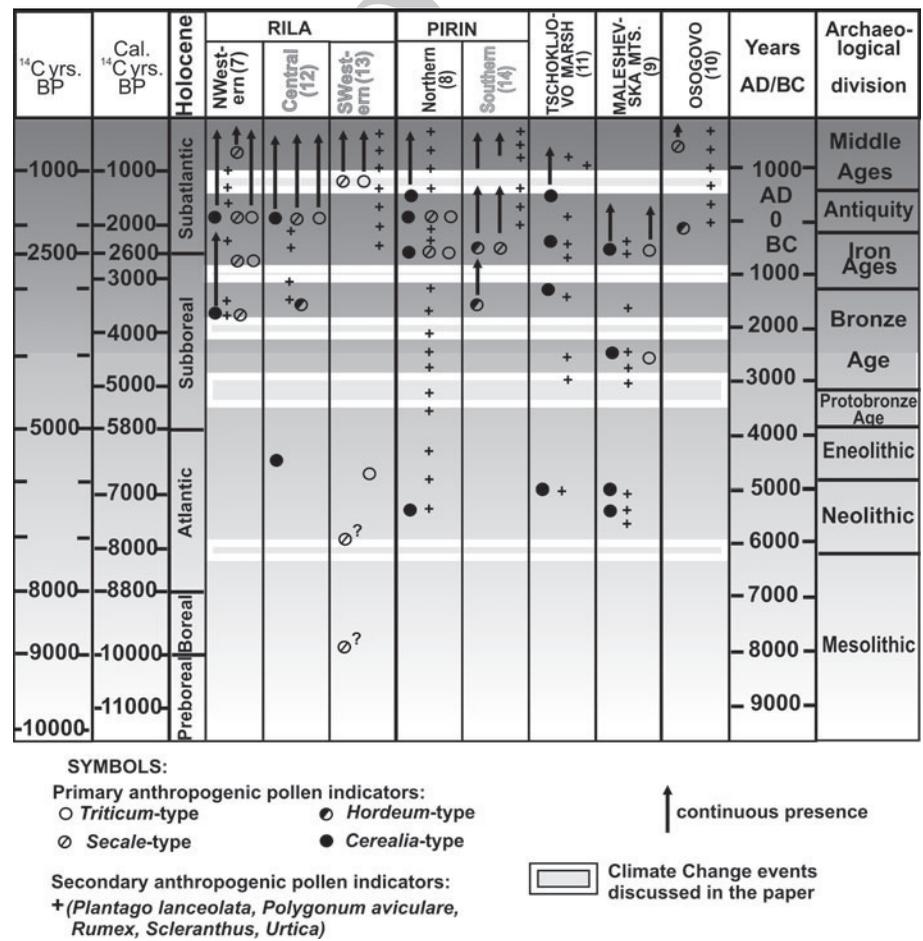


Fig. 4 Summary scheme for the anthropogenic indicators in the pollen diagrams of southwestern Bulgaria. The sites considered in detail in this paper are in *bold black* and further studies included in the summary, not further discussed here, are in *gray*. The numbers in *brackets* after each of the pollen records correspond to those given in Fig. 1. North Western Rila—Lake Trilistnika (7) ESM Fig. 1; Central Rila—Ostrezki Lakes (12) Tonkov and Marinova (2005); South Western Pirin—Dry Lake (13) Bozilova et al. (1986); Northern Pirin—Lake Ribno Banderishko (8) ESM Fig. 2; Southern Pirin—peat bog Mutorog (14) Panovska et al. (1995); Maleshevska Mountain—Peat bog Maleshevska (9) ESM Fig. 3; Osogovo Mountain—Peat bog Osogovo (10) ESM Fig 4; Konjavaska Mountain—Tshokljovo marsh (11) ESM Fig. 5



(2004) was used. No hiatuses were detected in any of the pollen records. The age-depth models for each site are given in the corresponding publication referred to above. In the most cases the radiocarbon age control of the records is based on dating of bulk sediment except for the profile from Lake Ribno Banderishko (Tonkov et al. 2002), where AMS-dating on terrestrial plant macrofossils was applied. The lithological units for each of the records are represented in the corresponding diagrams (ESM Figs 1-5). The pollen source area for these palynological archives corresponds to small-medium size basins. As the sites are situated in montane areas transportation of pollen upslope from lower elevations is also considered as shown by studies of surface moss samples from the region (Tonkov 2007).

Palynological indicators for the interpretation of human impact on the vegetation in the study area

In order to understand the nature and extent of human impact on the ancient vegetation and landscape, it is important to define the major indicators for anthropogenic impact. Based on the approach of Behre (1981), the pollen data information is evaluated with special emphasis on the anthropogenic impact on the landscape visible in the palaeoecological records.

The primary anthropogenic indicators of cereal cultivation such as Cerealia-type, *Triticum*-type, *Hordeum*-type, *Avena*-type and *Secale* are also used in the current study for estimation of changes in land use (Faegri and Iversen 1989). The methodological approach for the defining of pollen taxa as “secondary anthropogenic indicators” in the pollen diagrams from Central Europe, the Mediterranean region and Greece has been discussed in detail (Behre 1981, 1988; Bottema 1982; Bottema and Woldring 1990). To a considerable extent these ideas were also applied to the pollen diagrams from Bulgaria (Bozilova and Tonkov 1990; Huttunen et al. 1992; Filipovich and Stefanova 1998; Lazarova 1995; Marinova and Atanassova 2006; Tonkov et al. 2011).

Special attention needs to be given to Cerealia-type as the pollen of several wild-growing grass taxa also shows the morphological features of this type (Beug 2004), which means that it does not necessarily come from cultivated fields, but could have originated from the local wild-growing vegetation. Hence, the records of this type should be considered more cautiously and only in connection with the other indicators for anthropogenic change and considering both the local vegetation and the information on human occupation in the study area (for further discussion on the topic see Behre 2007, 2008). The same is true to great extent for *Hordeum*-type, *Avena*-type and *Secale* as wild-growing representatives of these types could also be found in southeastern Europe. The archaeobotanical evidence

from the Neolithic, Chalcolithic and Bronze Age period (from ca. 5900 to 1200 cal B.C.) shows that the main cereal crops of importance for the study area (Popova and Marinova 2007) were hulled wheats (*Triticum monococcum* and *T. dicoccum*) and barley (*Hordeum vulgare*). For this period *Triticum*-type could certainly be considered as a primary anthropogenic indicator, as the genus has no wild-growing representatives in the area under consideration. It is more difficult to consider *Hordeum*-type as a definite indicator of cultivation as several weedy and wild-growing species from the genus *Hordeum* occur in the study area. From the archaeobotanical information from the region of south-eastern Europe (Kroll 1991; Neef 2007; Popova 2009) it is known that cultivated rye (*Secale cereale*) gained importance from the Late Roman period onwards, but had occurred as a weed since the Late Iron Age (6th–4th century B.C.). Considering this, the *Secale* in the pollen diagrams found in sediments of earlier age than the Roman period is more likely to represent non-cultivated rye present as a weed or an element of the natural vegetation. Similarly for cultivated oats (*Avena sativa*) no evidence exists for its cultivation before the Roman age, but as a weed *Avena* sp. is known from the archaeobotanical record since the Late Bronze Age (Popova 2009). However the peaks of *Hordeum*-type, *Avena*-type and *Secale* should be attributed to anthropogenic activities, especially in connection with other evidence for cultivation and human impact.

Another pollen type indicative of anthropogenic activities is *Agrostemma githago*. However it can only really be used as an anthropogenic indicator from the Late Iron Age/Roman Age onwards when this plant becomes a constant element of the weed flora of the study area (Popova 2009). The regular appearance of *Juglans regia* in Mid-Holocene pollen records (Tonkov and Bozilova 1992b; Tonkov et al. 2002) from the study area indicated that this tree should be considered as native for this part of the country. Considering its more extensive occurrence after ca. 1400 cal. B.C. (corresponding to the Late Bronze Age) and the palynological evidence from adjacent regions (Bottema 1974, 2000; Jahns and Van den Bogaard 1998; Eastwood et al. 1998; Kaltenrieder et al. 2010) the pollen of *Juglans regia* could be interpreted as an indicator of anthropogenic activities from 1400 cal. B.C. onwards.

The secondary anthropogenic indicators characteristic of the mountains in southwestern Bulgaria include *Plantago lanceolata*-type, *Polygonum aviculare*-type, *Rumex*, *Scleranthus*, *Urtica*, *Cirsium*-type and *Juniperus*. They reflect permanent human occupation—ruderalization and stock-breeding. Special attention is paid to *Plantago lanceolata*-type, as a taxon that apart from human activity in settlement areas also reflects livestock grazing in the mountain meadows, but is also an important indicator of fallow land.

Quite often after forest clearing the pollen curves of Poaceae, *Artemisia*, Chenopodiaceae, *Plantago major/media*-type, *Urtica* and Cichorioideae (*Taraxacum*- or, according to Beug 2004, *Crepis*-type) increase and by careful evaluation of the dating these can also be used as secondary indicators for interpretation of human impact.

Some of the secondary anthropogenic indicators, for example *Artemisia* sp., Chenopodiaceae, Cichorioideae, *Cirsium*-type, *Plantago major/media*-type, *Rumex* etc. were also elements of the natural open vegetation during the Late Glacial and Early Holocene and for these periods they are not considered to indicate human impact. When occurring in subsequent periods these taxa are considered as anthropogenic indicators mainly only if they show parallel peaks in the pollen diagrams or appear in combination with other primary evidence of human impact on the vegetation.

Anthracological data

Anthracological analysis permits the reconstruction of the surrounding vegetation on areas where there are no sediments suitable for pollen analysis. In study areas where pollen bearing sediments near prehistoric sites have not yet been recovered, it allows the creation of a relatively good picture of the past vegetation and environments. Moreover, it represents the woodland used directly by the prehistoric population and allows the identification through wood charcoal macro remains of some taxa that could not be revealed by pollen analysis.

The anthracological data used for the current study comes from several archaeological sites from the study area (Fig. 1), all of them situated in the lowlands of the region between ca. 200 and 700 m a.s.l.

Together with the already published information on some of the sites (Marinova et al. 2002; Marinova and Thiebault 2008; Popova and Marinova 2007), additional information was used from samples from Kovacevo collected after 2003 and from further studies in the region (for the sites Balgarchevo, Slatina).

The laboratory study was carried out on charcoal (>2 mm) manually fractured in three anatomical planes (transversal, tangential and radial). The charcoal was examined with a reflected light microscope and counted. Specialized literature (Schweingruber 1990) and a reference collection of modern wood samples were used for the identification. Studies on the sample size required for adequate reconstruction of woodland vegetation in the archaeobotanical assemblages for the Western Mediterranean have shown that the best ca. 400–500 fragments per stratigraphic unit should be used (Chabal 1992). Thus for the sites in southeastern Europe, where the biodiversity is roughly comparable, at least 400 charcoal fragments per

stratigraphic layer are needed. In the currently studied material this requirement is met by four of the total of six sites shown in Fig. 3.

The counts of wood charcoal fragments identified for each site were made to calculate percentage proportions between the different vegetation types to which the charcoal taxa most probably belonged and are plotted in chronological order (Fig. 3).

The classification of the different charcoal taxa to vegetation types follows those already published and extensively discussed by Marinova and Thiebault (2008) and is based on the descriptions of the potential natural vegetation in the region given in Bondev (1991) and Bohn et al. (2000/2003). The type defined as “oak forests” includes the wood charcoal fragments identified as *Quercus* sp., *Acer* sp. and *Carpinus* sp. The group defined as “forest border/undergrowth” includes light-demanding taxa like *Cornus* sp., *Corylus avellana*, *Juniperus* sp., Prunoideae, Pomoideae and Rosaceae. The group “pine forests” includes the wood charcoal identified as *Pinus nigra/sylvestris* and coniferous, and finally taxa like *Alnus* sp., *Fraxinus* sp., *Salix/Populus* and *Ulmus* sp. are attributed to the group called “riparian woodland”. Being aware that some of these taxa could occur in more than one of the groups defined by us, the decision was made to assign them to the most probable of them.

Results

Palynological evidence

Several pollen diagrams from the mountains of southwestern Bulgaria with relatively good chronological control give information related to anthropogenic impact since the Neolithic.

The correlation of the vegetation development with the palynological data for the region is given in Fig. 2. In the following section the anthropogenic impact on the vegetation will be briefly presented against the background of the general vegetation development for the five selected palynological records.

Lake Trilistnika, Northwestern Rila Mts., ESM Fig. 1

The pollen data span the Late Glacial and the entire Holocene (Tonkov et al. 2008). The palynological record provides some information on the anthropogenic impact. This is not rich bearing in mind the high elevation (2,216 m a.s.l.) of the site studied, but is of importance in tracing human influence in the higher mountain area and on a regional scale.

One of the most important environmental changes for the Holocene occurred at the transition of pollen zones LT-4 and LT-5 (ca. 5900 cal. B.C. corresponding to the Early Neolithic) when the deciduous *Quercus*-forests with *Tilia*, *Ulmus* and *Corylus* declined. Coniferous forests dominated by *Pinus* (*P. sylvestris*, *P. peuce*) and *Abies alba* developed at higher altitudes. The decrease in *Betula* pollen indicates that in many places birch forests gave way to coniferous vegetation. The vertical migration of conifers like *A. alba* and *P. peuce* and the expansion of their areas was probably facilitated by an increase in precipitation and humidity (Davis et al. 2003), taking place after soils with humic horizons had developed (Bennett and Willis 1995). By this time the first sporadic records of *Plantago lanceolata*-type had occurred. From 1400 cal. B.C. (Late Bronze Age) the continuous presence of pollen of *Rumex* and *P. lanceolata*-type begins and the first pollen grains of *Juglans regia* appear. The establishment of *J. regia* pollen and the presence of cereal pollen (*Triticum*-type, *Secale*) points to an expansion of agriculture and cultivation of walnut in the foothills of the mountains, whereas *Rumex*, *P. lanceolata*-type and *Scleranthus* pollen indicate livestock-grazing in the mountain meadows.

Lake Ribno Banderishko, Northern Pirin Mts., ESM Fig. 2

The palynological data cover part of the Lateglacial and the entire Holocene (Tonkov et al. 2002). The site is situated at 2,190 m a.s.l. Comparable to the situation in the Rila Mts. at ca. 5900 cal. B.C. (Early Neolithic) important changes in the forest cover occurred. The mixed deciduous *Quercus*-forests with abundant *Tilia*, *Ulmus* and *Corylus* retreated and were replaced in many places by *Carpinus orientalis*/*Ostrya carpinifolia*, while further upslope communities of *C. betulus* developed. The favorable climatic and edaphic conditions triggered the formation of a coniferous belt dominated by *Pinus sylvestris*, *P. peuce* and *A. alba*. This is also the period when the first indications of human presence are recorded in the pollen diagram. The beginning of a continuous *P. lanceolata*-type pollen curve and the first appearance of Cerealia-type, *P. major/media*-type and *Polygonum aviculare*-type pollen is noteworthy. Later on, ca. 1000 cal. B.C., the uninterrupted presence of *Scleranthus* pollen begins. The most indicative signs of human impact were present after ca. 700 cal. B.C. with the continuous pollen curve of Cerealia-type and the regular presence of both *J. regia* and secondary anthropogenic pollen indicators.

Peat bog Maleshevska Mts., ESM Fig. 3

The palynological record from this lower mountain, situated in a western direction from the Struma river at about

1,700 m a.s.l., reveals the changes in the natural vegetation since 5500 cal. B.C. (Late Neolithic) (Tonkov and Bozilova 1992a). Three distinct periods are recognized in the pollen diagram.

During the earliest period (5500–2760 cal. B.C., Middle/Late Neolithic-Early Bronze Age) forests of *A. alba* with an admixture of *P. sylvestris* and *P. nigra* dominated above 1,000–1,200 m thus shaping the upper tree-line for more than the next 3,000 years. The wide distribution of the conifers was favored by an increase in precipitation and humidity during the Holocene climatic optimum. Below the coniferous belt, mixed oak forests (*Quercetum mixtum*) with abundant *Tilia*, *Ulmus* and *Corylus avellana* developed. The presence of Cerealia-type pollen observed for the last quarter of the 6th millennium B.C. might reflect increasing settlement activity and, related to this, cereal cultivation in the Struma valley.

During the next period (2760–800 cal. B.C., Early Bronze Age-Early Iron Age) a sharp decline in the areas occupied by *Abies* is observed around 2760 cal. B.C. Quite probably, this was caused by a decrease in humidity (roughly corresponding to Bond event 4) and was reinforced by the intensification of human disturbance in all vegetation belts. By ca. 800 cal. B.C. beech forests replaced the conifers, shaping the present tree-line. The uninterrupted presence of the anthropogenic pollen indicators (Cerealia-type, *P. lanceolata*-type, *Rumex*, *Scleranthus*) clearly indicates the practice of agriculture in the foothills of the mountain and extended sheep/goat and cattle-breeding in the lower mountain areas. The continuous presence of *Juglans* pollen dates back to the Early Bronze age.

The last period (800 cal. B.C.—present) is characterized by a strong anthropogenic influence on the vegetation. At lower elevations the oak forests were anthropogenically disturbed and replaced by secondary communities dominated by *Carpinus orientalis* and *Quercus pubescens*, while at higher elevations beech forests replaced the conifers.

Peat bog Osogovo Mts., ESM Fig. 4

The palynological data from this site provide a reliable basis for a palaeovegetation reconstruction for the last ca. 5000 years (Tonkov 2003). The site is situated at 1,720 m a.s.l. and the oldest pollen spectra reflect a characteristic vegetation pattern dominated by *Pinus* sp. and *Abies alba* which lasted until ca. 200 cal. B.C. The broad-leaved vegetation distributed at lower altitudes was composed of *Quercus* sp., *Carpinus betulus*, *Tilia* sp. and *Corylus avellana*. The first indication of human presence is the appearance of the pollen curves of *Rumex* and *P. lanceolata*-type at ca. 1500 cal. B.C. (transition to Late Bronze Age) synchronously with the decline of the conifers and the

578 rise of the pollen curve of *Fagus sylvatica*. In the course of
579 a thousand years beech has replaced the conifers and in this
580 process the anthropogenic impact should also be taken into
581 consideration. The high values of *Scleranthus* pollen found
582 throughout the entire profile are of local/extra-local origin
583 and are most probably indicative of stock-breeding prac-
584 ticed in the vicinity. The first peak of *Scleranthus* could be
585 assigned to the onset of the Late Bronze Age.

586 *Tschokljovo Marsh, Konjavka Mts., ESM Fig. 5*

587 The study site is nowadays a large marsh (870 m a.s.l.)
588 formed in a depression in the Konjavka Mts. (1,487 m)
589 near to the upper course of the Struma River. The pollen
590 data span the last 9,000 years (Tonkov and Bozilova
591 1992b; Bozilova and Tonkov 2007).

592 Indications of human presence are not recorded between
593 7000 and 5200 cal. B.C. when the mountain slopes sur-
594 rounding the marsh were covered by mixed oak forests
595 with *Tilia*, *Corylus avellana*, *Ulmus*, some *Carpinus ori-*
596 *entalis/Ostrya carpinifolia* and *C. betulus*. On the higher
597 mountain areas stands of *Pinus* sp. and *A. alba* occurred.

598 The following period of ca. 3,000 years (5200-ca.
599 2000 cal. B.C.) is characterized by a vast spread of *A. alba*
600 forests. The most interesting feature, however, is the con-
601 tinuous presence of *Juglans* pollen, probably indicating its
602 native distribution in this area rather than cultivation by the
603 local population.

604 During the subsequent period, which lasted until the
605 start of the Roman occupation (pollen zone transition T-4/
606 T-5), the palynological indications of human presence and
607 activity occur rather sporadically. The coniferous forests
608 composed of *A. alba* and *Pinus* were replaced by forests of
609 *F. sylvatica*, also around the start of the Roman period.

610 Wood charcoal analysis

611 The summary of the data from wood charcoal analysis is
612 given in Fig. 3. The dominant vegetation type around the
613 archaeological sites is deciduous oak forest and its repre-
614 sentatives dominate the anthracological assemblages. An
615 important characteristic of the wood charcoal assemblages
616 is the increasing proportion of light demanding shrubby
617 forest border and undergrowth elements of the oak forests.
618 At most of the sites this is shown by the rather high per-
619 centage proportions of *Cornus* sp. wood (Marinova and
620 Thiebault 2008; Marinova et al. 2002), followed by rep-
621 resentatives of Rosaceae, and to a lesser extent, of *Juni-*
622 *perus* sp. and *Corylus avellana*. Many of these small trees
623 and shrubs could not easily be detected by the palynolog-
624 ical investigations as they are not wind- but insect-pollin-
625 ated and/or were growing mainly in the lowlands, far from
626 the palynological sites available in the study area. The

anthracological records cover the period when farming 627
started in the study area and thus overcome some of the 628
above mentioned disadvantages of the palynological evi- 629
dence in tracing the first anthropogenic influence on the 630
vegetation. Figure 3 shows that, starting with the Early 631
Neolithic then until the Early Chalcolithic, the proportions 632
of the light-demanding woodland (given in Fig. 3 under 633
“Forest border/undergrowth”) and riparian woodland 634
increase in the anthracological assemblages. It can also be 635
seen that the Neolithic population certainly had access to 636
Pinus stands or woodland at lower elevations, most prob- 637
ably those were mainly *P. nigra* stands nowadays greatly 638
diminished by anthropogenic pressure (Bondev 1991). 639

Discussion 640

The palynological evidence on the anthropogenic changes 641
of the vegetation is summarized and correlated for the 642
different sub-regions in Fig. 4. This summary is based on 643
the palynological examples closely considered in this paper 644
(ESM Figs. 1-5) and on additional evidence described in 645
detail by Tonkov (2007). The development of the natural 646
landscapes is discussed in the context of the changing 647
human occupation in the area and of the climate change 648
events observed on a global scale discussed in Bond et al. 649
(2001), Mayewski et al. (2004) and Wanner et al. (2008). 650

First signs of human impact on the vegetation 651
(Neolithic to Early Bronze Age, 6200–2800 cal B.C.) 652

The up-to-date palynological evidence from southwestern 653
Bulgaria shows the response of the vegetation to the 8200 654
B.P. event, this response best manifested by the expansion of 655
Abies alba in the coniferous belts of Rila and Pirin moun- 656
tains. (Tonkov et al. 2008). The start of the Holocene climate 657
optimum facilitated the spread of mixed oak forests at low 658
and mid-altitudes, where later the first prehistoric settle- 659
ments were founded. The palynological records from the 660
area under consideration come from higher altitudes, in most 661
cases above 1,000 m, hence above the main area of activity 662
of the Neolithic population and of limited use for recon- 663
structing human impact on the vegetation for that period. For 664
the only site lower than this limit, Tschokljovo marsh 665
(870 m a.s.l.), the resolution of the palynological record is 666
not high for this period, and shows that oak forests with hazel 667
as a pioneering element were present in the lowlands. 668

The evidence from wood charcoal analysis is very useful 669
for tracing the woodland use in the initial stages of farming 670
introduction in the region, as this comes directly from the 671
prehistoric settlements. The anthracological as well as the 672
palynological records indicate wide distribution of decid- 673
uous forests dominated by *Quercus* sp. These were the 674

main woodland resource used by the Neolithic population. In the course of the Neolithic light-demanding trees and shrubs became more important in the oak woodlands. This change is related to the increased disturbance of the woodland by the local people, e.g. establishment of cultivating fields, grazing of animals, collecting fruits, fodder and firewood. This led to an increase of forest edge zones and secondary forests. Similar tendencies are also observed in the palynological records from Slovenia for the period of ca. 5500 cal. B.C., when no forest clearance occurred during the Neolithic period, but small-scale forest modifications, burning and coppicing were detected (Andric and Willis 2003). Moreover, the Neolithic land use strategies, involving coppicing and pollarding and forest pasture of small ruminants, favoured and enlarged such landscapes as is visible in the evidence from Central Europe (Kalis et al. 2003; Kreuz 2008; Bleicher and Herbig 2010; Gardner 2002; Magyari et al. in press). The evidence from Anatolia (Asouti and Hather 2001; Fairbairn et al. 2002; Riehl and Marinova 2008), the Balkans and Central Europe shows quite uniform land use strategies for broad parts of the continent during the Neolithic, resulting in slight modifications of the woodlands and leading to a more patchy character of the forests, with increasing diversity of their composition. Considering also the hypothesis of the uniformity of the Neolithic crop cultivation practices in Central Europe, South Eastern Europe and the Near East proposed by Bogaard (2004), a uniformity of general plant use strategies (including both field cultivation and land use) for this broad area could even be suggested. However palynological records with detailed chronological estimations as well as wood charcoal records concerning all stages of the Neolithic and the Chalcolithic are still needed to reconstruct reliably and more specifically the impact on the vegetation of the agriculture and land use for these periods which span over 1,800 years. Moreover the role of the fire in shaping the vegetation and landscape still needs to be explored by increasingly including micro- and macro-charcoal records, which at present are very scarcely available for the region.

In the pollen diagrams from Northern Pirin, Konjavska and Maleshevska mountains (ESM Figs. 2, 3, and 5) the first sporadic indications of cultivation of cereals (Cerealia-type) and grazing activities (*Cirsium*-type, *P. lanceolata*-type, *Polygonum aviculare*-type, *Rumex*, *Scleranthus*) appear during the Late Neolithic (ca. 5500–5000 cal. B.C.). This is related to more intensive Late Neolithic occupation of the region compared to the previous period. At the beginning of the Late Neolithic there are about 15 recorded settlements, during the second half there are 23, these being situated on the first or second terrace of the river, but higher up as well (Grębska-Kulowa and Kulow 2007). Hence, the first clearly noticeable anthropogenic indicators

in the palynological record coincide with the increasing number of settlements. It seems that the signal is rather of extra local character as it is visible in all of the palynological records considered (Fig. 4; ESM Figs. 1–5), including those from the higher mountain sites. The second half of the Late Neolithic is the period contemporary with the expansion of the deciduous and coniferous trees (formation of *Pinus* and *Abies* belt) in the mountains along the Struma valley. The spread of *A. alba* after 5250 cal. B.C. suggests an increase in humidity and temperature (Bozilova and Tonkov 1994). Against this background the indications of increasing light-demanding trees and of the enlarged area for obtaining firewood, present in the wood charcoal assemblages (Fig. 3) should also be interpreted as an indication of anthropogenically-driven opening of the landscape in the surroundings of the sites investigated.

The wood charcoal assemblages from the Early Chalcolithic (ca. 4900–4500 cal. B.C.) show a continued increase in the use of light-demanding and riparian woodland. The settlements decreased in number and moved to hilly areas with defendable locations (Grębska-Kulowa and Kulow 2007).

During the transition between the Late Chalcolithic to the start of the Early Bronze Age (ca. 3800–3200 cal. B.C.) there are almost no indications of anthropogenic impact on the vegetation, except a few weak peaks of secondary anthropogenic indicators such as *P. lanceolata*-type, *Scleranthus*, *Polygonum aviculare*-type and Cichorioideae. Such secondary anthropogenic indicators pointing to pasture activities are also present in the contemporary palynological records from south western Balkans (Sadori 2007). These are usually interpreted as a result of pasture activities. For this period the archaeological evidence also suggests subsistence relying on herding and the presence of fairly mobile groups (Leshtakov 2006).

Large scale human impact on the landscape traceable over the region (Late Bronze Age to Iron Age, 1400–100 cal. B.C.)

After a period with a shift to increasing pastoralism and less permanent settlements around 3200–2200 cal. B.C. the number of settlements in the study area during the Late Bronze Age (1400–1200 cal. B.C.) increased rapidly. For example, in the Middle Struma valley Grębska-Kulowa and Kulow (2007) report about 33 sites (compared with 3 sites for the preceding period). The Late Bronze Age is also the period with the next peak in the indicators of anthropogenic activities. This peak is much more pronounced than that in the Neolithic, and is shown in all of the palynological archives considered. Immediately after this peak the *F. sylvatica* curve starts to increase in the palynological record from the Maleshevska peat bog (ESM Fig. 3).

779 A similar situation with the rising of *Fagus* percentage
780 values after anthropogenic impact, and also with records of
781 fire activity was observed at the peat bog Beg Bunar in the
782 Osogovo mountains (Lazarova et al. 2009). This evidence
783 indicates clearly that at least in the lower mountain ranges
784 in the study area the spread of *F. sylvatica* was favored by
785 anthropogenic activities. This evidence is in accordance
786 with the estimations of Giesecke et al. (2007) for the
787 lowlands of central Europe and for the southern Prealps
788 (Valsecchi et al. 2008), where disturbance and anthropo-
789 genic clearance accelerated the growth of *F. sylvatica*
790 populations during the Holocene. In the higher mountains
791 (Lake Trilistnika, NW Rila; ESM Fig. 1) the appearance of
792 anthropogenic indicators such as *P. lanceolata*-type and
793 *Rumex* is recorded after 1400–1200 cal. B.C. (Late Bronze
794 Age) coinciding with the starting point of increase in
795 *F. sylvatica* and *Picea abies*. But it should be mentioned
796 that this is also a period when an increase of moisture
797 availability favourable for both tree species was recon-
798 structed for the Rila Mountains (Tonkov and Marinova
799 2005), so climatically driven increase of the *F. sylvatica*
800 population is quite plausible for those higher elevations.
801 However more multi-proxy records with finer chronologi-
802 cal resolution from the region are needed to answer this
803 question reliably.

804 For the subalpine zones convincing evidence of human
805 presence is observed in the Rila Mountains in the pollen and
806 plant macrofossil diagrams from the Ostrezki Lakes (Ton-
807 kov and Marinova 2005). The pollen curves of taxa such as
808 Cerealia-type, *P. lanceolata*-type, *Rumex*, *Scleranthus* and
809 *Urtica*, along with macro-charcoals, appeared at 1770 cal.
810 B.C., while numerous charcoal fragments indicative of forest
811 fires were found from later, during ca. 800 cal. B.C. After ca.
812 900–800 cal. B.C. the palynological records from the
813 northern Pirin (this paper, and also Stefanova et al. 2003),
814 Maleshevska and Osogovo mountains also show pro-
815 nounced peaks of anthropogenic indicators indicating
816 mainly pasture (e.g. *Artemisia*, *Cirsium*-type, *P. lanceolata*,
817 *Rumex*, *Scleranthus*) and a partial deforestation of the area.
818 At many places the upper tree-line was artificially lowered
819 in order to extend the high mountain pasture land. At lower
820 elevations the mixed oak forests were destroyed and com-
821 munities of *Carpinus orientalis* and *Quercus pubescens*,
822 secondary in origin, were formed. The onset of this period
823 with prominent anthropogenic activities is contemporane-
824 ous with the Early Iron Age in the area. It is also marked by
825 the end of a dry climatic interval from approx. 1100 to
826 900 cal. B.C. also registered in the Hungarian and Thracian
827 plains (Chapman et al. 2009).

828 The period of ca. 800 B.C. onwards for the Eastern
829 Mediterranean is characterized by an increase in humidity
830 and temperatures (Bar-Matthews et al. 1999; Lamy et al.
831 2006). Since 800 B.C. and especially around 500–400

832 B.C. human impact on the vegetation becomes clearly
833 pronounced and continuous on a large scale. This is also
834 clearly visible in the adjacent areas (Andric 2002; Atha-
835 nasiadis et al. 2000; Eastwood et al. 1998; Feurdean et al.
836 2010; Jahns 1993; Knipping et al. 2008; Lazarova et al.
837 2011; Mudie et al. 2007; Triantaphyllou et al. 2010). These
838 changes, except for the cultivation of olive which cannot
839 thrive in the study area, strongly resemble the Beyşehir
840 occupation phase (for more recent discussion on the topic
841 see Bakker et al. 2011). The current evidence confirms the
842 strong element of continuity in land use between the Late
843 Bronze Age settlers and the Early Iron Age communities
844 and the increasing intensity of forest clearance during the
845 Iron Age proposed for the region of southeastern Europe by
846 Chapman et al. (2009).

Roman and medieval landscape management 847
and land use 848

The trends observed during the Iron Age (800 B.C.–A.D. 100) 849
continued into the Roman period (from ca. A.D. 100–400). 850
The evidence for intensive deforestation and spread of 851
secondary xerothermic vegetation (increases or peaks in 852
Juniperus, *Carpinus orientalis/Ostrya carpinifolia*, 853
Scleranthus, Cichorioideae, *Cirsium*-type, *Artemisia* etc.) 854
related to stock-breeding as well as to field and tree culti- 855
vation is very well manifested in the pollen diagrams under 856
consideration. In most of the lower mountains in the study 857
area, including the records from Maleshevska, Osogovo and 858
Tshokljovo (ESM Figs. 3, 4 and 5) considered here, *Abies* 859
alba played a significant role until the Hellenistic/Roman 860
period. Evaluating the Holocene palaeoecological record of 861
A. alba in the study area, Bozilova and Tonkov (1994) 862
consider this reduction to be mainly anthropogenically 863
driven. Similar tendencies for anthropogenically influenced 864
reduction of *A. alba* populations over the last 2,000 years 865
are also observed in North Western Romania (Feurdean and 866
Willis 2008) and in Slovenia (Andric 2002). Arboriculture, 867
a typical element of Roman land use, is only partly traceable 868
by the increase of the *Juglans* curves and in some of the 869
records considered the curves of *Vitis sylvestris* and *Cas- 870
tanea sativa* also rise slightly prior to the Roman period. 871
The finds of *Juglans regia* are also very common and fre- 872
quent in the archaeobotanical record of the region for the 873
Iron Age and especially the Roman period (Popova 2009). 874
In general, during the Roman period the local agricultural 875
economy flourished as the Romans introduced a broader 876
spectrum of crops, advanced methods of field cultivation 877
and arboriculture (Bozilova et al. 1994). 878

879 The palaeoecological records considered here provide
880 little information on medieval land use as most of them end
881 before the start of the Middle Ages, or have no radiocarbon
882 age control for this period. However, further analyses and

more precise palaeoecological archives for this period are needed to elucidate the medieval human impact on the vegetation and the landscapes in the study area. The only evidence from this time is preserved in the Osogovo peat bog. There, the continuous presence of *Secale* and *Juniperus* pollen since ca. cal. A.D. 1200 indicates the rye cultivation typical of the period and increasing deforestation of the mountain. Here the strong and continuous presence of anthropogenic indicators coincides with an increase in *Fagus* woodlands.

Conclusions

1. The first major human impact, visible on a regional scale in the palynological records, appeared in the final stages of the Neolithic (ca. 5000 cal B.C.), when very sporadic indications of cereal cultivation and grazing occurred.
2. Anthropogenic vegetation change began with a slight modification of the xerothermic oak forests in the lowlands, the main area of activity of the Neolithic population, and with the stimulation of the development of light-demanding and in many cases fruit-producing trees and shrubs like *Cornus* sp., *Corylus avellana*, Maloideae, Prunoideae, *Rosa* sp. etc. The evidence for the Neolithic land use practices is in accordance with that from Anatolia and Central Europe.
3. The first signs of large scale deforestation and more extensive land use are recorded during the Late Bronze Age (ca. 1450–1200 cal. B.C.). In the lower mountains they are followed at many places by the continuous spread of *F. sylvatica* pointing to the favourable influence of the anthropogenic activities on its expansion. They are also marked by the first fairly continuous presence of indicators of pasture and fire activities in the mountainous areas.
4. The continuous presence of *Juglans* pollen from the Early Bronze Age and especially the Late Bronze Age points to its cultivation starting from this period. Its sporadic occurrence in the earlier periods and low percentage values suggest natural occurrence in the area.
5. The next peak of the anthropogenic impact on the vegetation and landscape occurred from 800 cal. B.C. onwards. It reached its maximum during the Roman period when annual crop and tree cultivation, deforestation and selective timber harvesting (*A. alba*) reached a vast extension.
6. The current study shows mainly the general tendencies in the human impact on vegetation and landscape that could be inferred from palaeobotanical data. Many questions on the exact timing and extend of those

impacts still remain and only future research focusing on high resolution records with more extended age control will help to gain a detailed and comprehensive picture of the human landscapes formed under the interaction of climate and cultural change.

Acknowledgments This paper is a contribution to Project N DID 02/26/2009 *Ecological Crises in Bulgaria during the Holocene—VII–III millennium B.C.* supported by the National Science Fund, Ministry of Education, Youth and Science in Sofia. The first author would like also to thank Bogdan Atanassov and Johan Bakker for valuable comments and discussion on earlier versions of the manuscript. The authors would like also to thank the guest editor Wiebke Kirleis and two anonymous reviewers for the valuable comments and suggestions which helped to improve the paper.

References

- Andric M (2002) The Holocene vegetation dynamics and the formation of Neolithic to Present-day Slovenian landscape. *Doc Praehistorica* 28:133–175
- Andric M, Willis K (2003) The phytogeographical regions of Slovenia: a consequence of natural environmental variation or prehistoric human activity? *J Ecol* 91:807–821
- Asouti E, Hather J (2001) Charcoal analysis and the reconstruction of ancient woodland vegetation in the Konya Basin, south-central Anatolia, Turkey: results from the Neolithic site Catalhöyük East. *Veget Hist Archaeobot* 10:23–32
- Athanasiadis N, Tonkov S, Atanassova J, Bozilova E (2000) Palynological study of Holocene sediments from Lake Doirani in northern Greece. *J Paleolimnol* 24:331–342
- Atanassov B, Petkov V, Kulow I, Grębska-Kulowa M, Dumanov B, Uzunov Z, Garbov D, Chukalev K, Gercheva D (2010) Expedition Struma during 2010 [Ekspedicia Struma prez 2010]. In: Gergova D (ed) *Archaeological findings and excavations 2009*, pp 676–679
- Bakker J, Paulissen E, Kaniewski D, De Laet V, Verstraeten G, Waelkens M (2011) Man, vegetation and climate during the Holocene in the Territory of Sagalassos, Western Taurus mountains, SW Turkey. *Veget Hist Archaeobot*. doi:10.1007/s00334-011-0312-4 (this volume)
- Bar-Matthews M, Ayalon A, Kaufman A, Wasserburg GJ (1999) The Eastern Mediterranean paleoclimate as a reflection of regional events: Soreq cave, Israel. *Earth Planet Sci Lett* 166:85–95
- Behre K-E (1981) The interpretation of anthropogenic indicators in pollen diagrams. *Pollen Spores* 23:225–245
- Behre K-E (1988) The role of man in European vegetation history. In: Huntley B, Webb T III (eds) *Vegetation history, handbook of vegetation science*. Kluwer, Dordrecht, pp 633–672
- Behre K-E (2007) Evidence for Mesolithic agriculture in and around central Europe? *Veget Hist Archaeobot* 16:203–219
- Behre K-E (2008) Comment on: Mesolithic agriculture in Switzerland? A critical review of the evidence by W. Tinner, E. H. Nielsen and A. F. Lotter. *Quat Sci Rev* 27:1,467–1,468
- Bennett K, Willis K (1995) The role of ecological factors in controlling vegetation dynamics on long temporal scale. *Giornale Botanico Italiano* 129:243–354
- Beug H-J (2004) *Leitfaden der Pollenbestimmung für Mitteleuropa und angrenzende Gebiete*. Pfeil, München
- Bleicher N, Herbig CH (2010) Der Federsee: Landschaft und Dynamik im Neolithikum. In: Matuschik I, Strahm C,

- 992 Eberschweiler B, Fingerlin G, Hafner A, Kinsky M, Mainberger
993 M (eds) Vernetzungen. Aspekte Siedlungsarchäologischer Forsch-
994 chung. Festschrift für Helmut Schlichtherle. Lavori, Freiburg,
995 pp 95–112
- 996 Bogaard A (2004) Neolithic Farming in central Europe: an archae-
997 obotanical study of crop husbandry practices. Routledge,
998 Oxon
- 999 Bohn U, Gollub G, Hettwer C, Neuheuslova Z, Schlüter H, Weber H
1000 (2000/2003) Map of the natural vegetation of Europe. Scale:
1001 1:2,500,000, Parts 1, 2 and 3. Landwirtschaftsverlag, Bonn
- 1002 Bond G, Kromer B, Beer J, Muscheler R, Evans MN, Showers W,
1003 Hoffmann S, Lotti-Bond R, Hajdas I, Bonani G (2001) Persistent
1004 solar influence on north Atlantic climate during the Holocene.
1005 Science 294:2,130–2136
- 1006 Bondev I (1991) The vegetation of Bulgaria. Map 1:600,000 with
1007 explanatory text. St. Kliment Ohridski University Press, Sofia (in
1008 Bulgarian with English summary)
- 1009 Borojevic K (2006) Terra and silva in the Pannonian Plain: Opovo
1010 Agro-Gathering in the Late Neolithic. BAR Internat Ser 1563.
1011 Archaeopress, Oxford
- 1012 Bottema AS (1974) Late Quaternary vegetation history of northwest-
1013 ern Greece. Thesis. Rijksuniversiteit Groningen
- 1014 Bottema AS (1982) Palynological investigations in Greece with
1015 special reference to pollen as an indicator of human activity.
1016 Palaeohistoria 24:257–289
- 1017 Bottema AS (2000) The Holocene history of walnut, sweet-chestnut,
1018 manna-ash and plane tree in the Eastern Mediterranean. PAL-
1019 LAS 52:35–59
- 1020 Bottema AS, Woldring H (1990) Anthropogenic indicators in the
1021 pollen record of the Eastern Mediterranean. In: Bottema AS,
1022 Entjes-Nieborg G, Van Zeist W (eds) Man's role in the shaping
1023 of the Eastern Mediterranean landscape. Balkema, Rotterdam,
1024 pp 231–264
- 1025 Boyadzhiev Y (2007) Absolute Chronology of the Neolithic and
1026 Eneolithic Cultures in the Valley of Struma. In: Todorova H,
1027 Stefanovic M, Ivanov G (eds) The Struma/Strymon river valley
1028 in prehistory. In the Steps of James Harvey Gaul, volume 2,
1029 Sofia, pp 309–316
- 1030 Boyadzhiev Y (2009) Early Neolithic cultures on the territory of
1031 Bulgaria. In: Gatsov I, Boyadzhiev Y (eds) The First Neolithic
1032 Sites in Central/South-East European transect volume I: Early
1033 Neolithic Sites on the territory of Bulgaria. BAR International
1034 Ser 2048, pp 7–44
- 1035 Bozilova E, Tonkov S (1990) The impact of man on the natural
1036 vegetation in Bulgaria from the Neolithic to the middle ages. In:
1037 Bottema S, Entjes-Nieborg G, Van Zeist W (eds) Man's role in
1038 the shaping of the Eastern Mediterranean landscape. Balkema,
1039 Rotterdam, pp 327–332
- 1040 Bozilova E, Tonkov S (1994) The postglacial distribution patterns of
1041 Abies in Bulgaria. In: Lotter A, Ammann B (eds) Festschrift
1042 Gerhard Lang. Beiträge zur Systematik und Evolution, Floristik
1043 und Geobotanik, Vegetationsgeschichte und Paläoökologie. Diss
1044 Bot 234, Cramer, Berlin Stuttgart, pp 215–223
- 1045 Bozilova E, Tonkov S (2007) Palaeoecological evidence of the main
1046 postglacial vegetation and climate changes in southwestern
1047 Bulgaria from the neolithic to modern times. In: Todorova H,
1048 Stefanovic M, Ivanov G (eds) The Struma/Strymon River valley
1049 in prehistory. In the Steps of James Harvey Gaul, vol 2, Sofia,
1050 pp 531–534
- 1051 Bozilova E, Tonkov S, Pavlova D (1986) Pollen and plant macrofossil
1052 analyses of the Lake Sucho Ezero in the south Rila mountains.
1053 Annu Sofia Univ Fac Biol 80:48–57
- 1054 Bozilova E, Tonkov S, Popova TZ (1994) Forest clearance, land use
1055 and human occupation during the Roman colonization in
1056 Bulgaria. In: Frenzel B (ed) Palaeoclimate research 10. Fischer,
1057 Stuttgart, pp 37–44
- Chabal L (1992) La représentativité paléo-écologique des charbons de
bois archéologiques issus du bois de feu. Bulletin de la Société
Botanique de France 139:213–236
- Chapman J, Magyari E, Gaydarska B (2009) Contrasting subsistence
strategies in the early Iron Age? New Results from the Alföld
plain, Hungary and the Thracian plain, Bulgaria. Oxf J Archaeol
28:155–187
- Chohadzhiev S (2007) Neolithni I halkolitni kulturi v basejna na reka
struma [Neolithic and Chalcolithic cultures in the Struma river
basin], Veliko Tarnovo University Press, Veliko Tarnovo (in
Bulgarian)
- Davis B, Brewer S, Stevenson A, Guiot J, Data Contributors (2003)
The temperature of Europe during the Holocene reconstructed
from pollen data. Quat Sci Rev 22:1,701–1,716
- Delev P (2002) The middle Mesta region in antiquity. In: Bozkova A,
Delev P (eds) Koprivlen vol 1: rescue archaeological investiga-
tions along the Gotse Delchev—Drama Road 1998–1999. NOUS
Publishers, Sofia, pp 13–28
- Eastwood WJ, Roberts N, Lamb HF (1998) Palaeoecological and
archaeological evidence for human occupation in southwest
Turkey: the Beyşehir occupation phase. Anatol Stud 48:69–86
- Faegri K, Iversen J (1989) Textbook of pollen analysis. In: Faegri K,
Kaland PE, Krzywinski K (eds) 4th edn. Wiley, Chichester
- Fairbairn A, Asouti E, Near J, Martinoli D (2002) Macro-botanical
evidence for plant use at Neolithic Çatalhöyük, south-central
Anatolia, Turkey. Veget Hist Archaeobot 11:41–54
- Feurdean A, Willis KJ (2008) Long-term variability of *Abies alba* in
NW Romania: implications for its conservation management.
Divers Distrib 14:1,004–1,017
- Feurdean A, Willis KJ, Parr C, Tantau I, Farcas S (2010) Postglacial
patterns in vegetation dynamics in Romania: homogenization or
differentiation? J Biogeogr 37:2,197–2,208
- Filipovich L, Stefanova I (1998) Anthropogenic changes in the
vegetation of the Balkan Range according to data obtained from
pollen and macrofossil analyses. Phytologia Balcanica 4:37–44
- Gaillard M-J (2007) Pollen methods and studies: archaeological
applications. In: Elias S (ed) Encyclopedia of quaternary science.
Elsevier, Amsterdam, pp 2,571–2,595
- Gardner A (2002) Neolithic to Copper Age woodland impacts in
northeast Hungary? Evidence from the pollen and sediment
chemistry records. Holocene 12:453–541
- Gatsov I, Boyadzhiev Y (2009) The First Neolithic Sites in Central/
South-East European transect volume I: Early Neolithic Sites on
the territory of Bulgaria. BAR Internat Ser 2048, Archaeopress,
Oxford
- Giesecke T, Hickler T, Kunkel T, Sykes MT, Bradshaw RHW (2007)
Towards an understanding of the Holocene distribution of *Fagus
sylvatica* L. J Biogeogr 34:118–131
- Görsdorf J, Bojadziev J (1996) Zur absoluten Chronologie der
bulgarischen Urgeschichte. Berliner 14C-Datierungen von bul-
garischen archäologischen Fundplätzen. Eurasia Antiqua 2:
105–173
- Grębska-Kulova M (2005) Cultural changes in the second half of the
6th mill. b.c. in Southwestern Bulgaria. In: Nikolov V, Bačvarov
K, Kalchev P (eds) Praehistoric thrace. Proceedings of the
international symposium, Stara Zagora, 30.09.-04.10.2003. Bulg
Acad of Sci, Sofia, pp 133–145
- Grębska-Kulowa M, Kulow I (2007) Prehistorical Sites in the Middle
Struma River Valley between the end of the VIIth mill b.c. and
the Beginning of the Vth mill. b.c. In: Stefanovic M, Todorova H
(eds) The Struma/Strymon River Valley in Prehistory. In the
Steps of James Harvey Gaul, vol 2, Sofia, pp 279–296
- Huttunen A, Huttunen R-L, Vasary Y, Panovska H, Bozilova E
(1992) Late glacial and Holocene history of flora and vegetation
in the Western Rhodopes Mountains, Bulgaria. Acta Bot Fennica
144:63–80

- Jahns S (1993) On the Holocene vegetation history of the Argive Plain (Peloponnese, southern Greece). *Veget Hist Archaeobot* 2:187–203
- Jahns S, Van den Bogaard C (1998) New palynological and tephrostratigraphical investigations of two salt lagoons on the island of Mljet, south Dalmatia, Croatia. *Veget Hist Archaeobot* 7:219–234
- Kalis AJ, Merkt J, Wunderlich J (2003) Environmental changes during the Holocene climatic optimum in central Europe—human impact and natural causes. *Quat Sci Rev* 22:33–79
- Kaltenrieder P, Procacci G, Vannièrè B, Tinner W (2010) Vegetation and fire history of the Euganean Hills (Colli Euganei) as recorded by Lateglacial and Holocene sedimentary series from Lago della Costa (northeastern Italy). *Holocene* 20:679–695
- Knipping M, Müllenhof M, Brückner H (2008) Human induced landscape changes around Bafa Gölü (western Turkey). *Veget Hist Archaeobot* 17:365–380
- Kopravev IE (2002) Geography of Bulgaria. Physical geography. Socio-economic geography. ForComPublishers, Sofia
- Kreuz A (2008) Closed forest or open woodland as natural vegetation in the surroundings of Linearbandkeramik settlements? *Veget Hist Archaeobot* 17:51–64
- Kreuz A, Marinova E, Schäfer E, Wiethold J (2005) A comparison of Early Neolithic crop and weed assemblages from the Linearbandkeramik and the Bulgarian Neolithic cultures: differences and similarities. *Veget Hist Archaeobot* 14:237–258
- Kroll H (1991) Südosteuropa. In: Van Zeist W, Wasylikowa K, Behre K-E (eds) Progress in old world palaeoethnobotany. Balkema, Rotterdam, pp 161–178
- Lamy F, Arz HW, Bond GC, Bahr A, Pätzold J (2006) Multicentennial-scale hydrological changes in the Black Sea and northern Red Sea during the Holocene and the Arctic/North Atlantic Oscillation. *Paleoceanography* 21, PA1008, doi:10.1029/2005PA001184
- Lazarova M (1995) Human impact on the natural vegetation in the region of Lake Srebarna and the Mire Garvan (Northeastern Bulgaria)—palynological and paleoethnobotanical evidence. In: Bozhilova E, Tonkov S (eds) Advances in Holocene palynology in Bulgaria. Pensoft Publishers, Sofia-Moskow, pp 47–67
- Lazarova M, Tonkov S, Snowball I, Marinova E (2009) 6. Peat-bog Begbunar (Osogovo Mountains, south-west Bulgaria): four millennia of vegetation history. *Grana* 48:147–179
- Lazarova M, Koutsios A, Kontopoulos N (2011) Holocene vegetation history of the Kothi lagoon (northwest Peloponnesus, Greece). *Quat Int*. doi:10.1016/j.quaint.2009.10.036
- Leshtakov K (2006) The Bronze Age in the Upper Thracian plain [Bronzovata epoha v gornotrakijskata hizina]. *Annuary of the Sofia University—Archaeology* vol 3, pp 141–216 (in Bulgarian with English summary)
- Lichardus-Ippen M, Demoule J-P, Perničeva L, Grębska-Kulova M (2006) Kovachevo, an early Neolithic site in South-West Bulgaria and its importance for the European Neolithization. In: Gatsov I, Schwarzberg H (eds) Aegean-Marmara-Black Sea: the present state of research on the early Neolithic. *Schriften des Zentrums für Archäologie und Kulturgeschichte des Schwarzmeerraumes* 5, pp 83–94
- Magyari E, Chapman E, Francis M, De Guzman M (in press) Neolithic human impact on the landscapes of North-East Hungary inferred from pollen and settlement records. *Veget Hist Archaeobot*
- Marinova E (2006) Vergleichende palaeoethnobotanische Untersuchung zur Vegetationsgeschichte und zur Entwicklung der prähistorischen Landnutzung in Bulgarien. *Diss Bot* 401. Cramer, Stuttgart, pp 1–164
- Marinova E, Atanassova J (2006) Anthropogenic impact on vegetation and environment during the Bronze Age in the area of Lake Durankulak, NE Bulgaria: pollen, microscopic charcoal, non-pollen palynomorphs and plant macrofossils. *Rev Palaeobot Palynol* 141:165–178
- Marinova E, Popova TZ (2008) *Cicer arietinum* (chick pea) in the Neolithic and Chalcolithic of Bulgaria: implications for cultural contacts with the neighbouring regions? *Veget Hist Archaeobot* 17(Suppl 1):73–80
- Marinova E, Thiebault S (2008) Anthracological analysis from Kovacevo, southwest Bulgaria: woodland vegetation and its use during the earliest stages of the European Neolithic. *Veget Hist Archaeobot* 17:223–231
- Marinova E, Tchakalova E, Stoyanova D, Grozeva S, Dočeva E (2002) Ergebnisse archäobotanischer Untersuchungen aus dem Neolithikum und Chalcolithikum in Südwestbulgarien. *Archaeol Bulgaria* 6:1–11
- Mayewski P, Rohling E, Stager C, Karlen W, Maascha K, Meeker D, Meyerson E, Gasse F, Van Kreveld S, Holmgren K, Lee-Thorp J, Rosqvist G, Rack F, Staubwasser M, Schneider R, Steig E (2004) Holocene climate variability. *Quat Res* 62:243–255
- Mudie PJ, Marret F, Aksu AE, Hiscott RN, Gillespie H (2007) Palynological evidence for climatic change, anthropogenic activity and outflow of Black Sea water during the late Pleistocene and Holocene: Centennial- to decadal-scale records from the Black and Marmara Seas. *Quat Int* 167–168:73–90
- Neef R (2007) Archäobotanische Untersuchungen im spätantiken Iatrus/Krivina (Grabungskampagnen 1992–2000). In: Von Bülow G, Böttger B, Conrad S, Döhle B, Gomolka-Fuchs G, Schönert-Geiss E, Stancev D, Wachtel K (eds) Iatrus-Krivina. Spätantike Befestigung und frühmittelalterliche Siedlung an der unteren Donau. *Limesforschungen* 28. Philipp von Zabern, Mainz, pp 415–445
- Nikolov V (2007) Problems of the early stages of the Neolithisation in the Southeast Balkans. In: Spataro M, Biagi P (eds) A short walk through the Balkans: the first farmers of the carpathian basin and adjacent regions. *Società Preistoria Protostoria Friuli-V.G. Trieste Quaderno* 12:183–188
- Panovska H, Bozilova E, Tonkov S (1995) A palaeoecological investigation on the vegetation history in the Southern Pirin Mts. (Southwestern Bulgaria). In: Bozilova E, Tonkov S (eds) Advances in Holocene palaeoecology in Bulgaria. Pensoft, Sofia, pp 32–46
- Perničeva L (1995) Prehistoric cultures in the middle Struma Valley: Neolithic and Eneolithic. In: Bailey D, Panayotov I (eds) Prehistory of Bulgaria. *Monographs in World Archaeology* 22, I. Madison, Wisconsin, pp 99–128
- Popova T (2009) Paleobotanic catalogue of the studied sites and studied remains (debries) in the territory of Bulgaria (1980–2008). *Interdiscip Stud* 20–21:71–165
- Popova T, Bozhilova E (1997) Palaeoecological and palaeoethnobotanical data for the Bronze Age in Bulgaria. In: Stefanovich M, Henrieta T, Hauptmann H (eds) In the steps of James Harvey Gaul, vol 1. Sofia, pp 391–397
- Popova T, Marinova E (2007) Palaeoethnobotanical data in South-Western Region of Bulgaria In: Todorova H, Stefanovic M, Ivanov G (eds) The Struma/Strymon river valley in prehistory. In the Steps of James Harvey Gaul, vol 2. Sofia, pp 523–532
- Riehl S, Marinova E (2008) Mid-Holocene vegetation change in the Troad (W Anatolia): man-made or natural? *Veget Hist Archaeobot* 17:297–312
- Roberts N, Jones MD, Benkaddour A, Eastwood WJ, Filippi ML, Frogley MR, Lamb HF, Leng MJ, Reed JM, Stein M, Stevens L, Valero-Garcés B, Zanchetta G (2008) Stable isotope records of Late Quaternary climate and hydrology from Mediterranean lakes: the ISOMED synthesis. *Quat Scie Rev* 27(25–26): 2,426–2,441
- Roberts N, Eastwood WJ, Kuzucuoğlu C, Fiorentino G, Caracuta V (2011) Climatic, vegetation and cultural change in the eastern

- 1256 Mediterranean during the mid-Holocene environmental transi- 1298
 1257 tion. *Holocene* 21:147–162 1299
 1258 Sadori L (2007) Pollen records, postglacial. Southern Europe. 1300
 1259 Encyclopaedia of quaternary sciences. Elsevier, Amsterdam, 1301
 1260 pp 2,763–2,773 1302
 1261 Schweingruber FH (1990) Anatomie europäischer Hölzer. Ein 1303
 1262 Atlas zur Bestimmung europäischer Baum-, Strauch- und 1304
 1263 Zwergstrauchhölzer. Haupt, Bern Stuttgart 1305
 1264 Stefanova I, Ognjanova-Rumenova N, Hofmann W, Ammann B 1306
 1265 (2003) Late Glacial and Holocene environmental history of the 1307
 1266 Pirin Mountains (SW Bulgaria): a paleolimnological study of 1308
 1267 Lake Dalgoto (2,310 m). *J Paleolimnol* 30:95–111 1309
 1268 Stefanovich M, Bankoff A (1998) Kamenska Čuka 1993–1995. In: 1310
 1269 Stefanovich M, Todorova H, Hauptman H (eds) In the steps of 1311
 1270 James Harvey Gaul, vol 1, Sofia, pp 255–338 1312
 1271 Todorova H, Vaisov I (1989) Das Neolithikum in Bulgarien. Nauka i 1313
 1272 Izkustvo, Sofia (in Bulgarian with German summary) 1314
 1273 Tonkov S (1988) Sedimentation and local vegetation development 1315
 1274 of a reference site in SW Bulgaria. In: Lang G, Schluchter C 1316
 1275 (eds) Lake, mire and river environments. Balkema, Rotterdam, 1317
 1276 pp 99–101 1318
 1277 Tonkov S (2003) A 5000-year pollen record from Osogovo Mountains, 1319
 1278 Southwestern Bulgaria. In: Tonkov S (ed) Aspects of palynology 1320
 1279 and palaeoecology. Pensoft, Sofia-Moscow, pp 233–244 1321
 1280 Tonkov S (2007) Paleobotanical and paleoecological investigation on 1322
 1281 the postglacial vegetation history in southwestern Bulgaria. DSc 1323
 1282 Thesis, Sofia (in Bulgarian) 1324
 1283 Tonkov S, Bozilova E (1992a) Pollen analysis of peat-bog in 1325
 1284 Maleshevska Mountain (Southwestern Bulgaria). *Ann Sofia* 1326
 1285 *Univ Fac Biol* 83:11–22 (in Bulgarian with English summary) 1327
 1286 Tonkov S, Bozilova E (1992b) Paleocological investigation of 1328
 1287 Tschokljovo marsh (Konjavaska Mountain). *Ann Sofia Univ Fac* 1329
 1288 *Biol* 83:5–16 1330
 1289 Tonkov S, Marinova E (2005) Pollen and plant macrofossil analyses 1331
 1290 of mid-Holocene radiocarbon dated profiles from two subalpine 1332
 1291 lakes in Rila Mountains, Bulgaria. *Holocene* 15:663–671 1333
 1292 Tonkov S, Panovska H, Possnert G, Bozilova E (2002) The Holocene 1334
 1293 vegetation history in the Northern Pirin Mountain, southwestern 1335
 1294 Bulgaria: pollen analysis and radiocarbon dating of core from 1336
 1295 Lake Ribno Banderishko. *Holocene* 12:201–210 1337
 1296 Tonkov S, Bozilova E, Possnert G, Velcev A (2008) A contribution to 1338
 1297 the postglacial vegetation history of the Rila Mountains, 1339
 Bulgaria: the pollen record of Lake Trilistnika. *Quat Intern* 1298
 190:58–70 1299
 Tonkov S, Beug H-J, Bozilova E, Filipova-Marinova M, Jungner H 1300
 (2011) Palaeoecological studies at the Kaliakra area, northeast- 1301
 ern Bulgarian Black Sea coast: 6000 years of natural and 1302
 anthropogenic change. *Veg Hist Archaeobot* 20:29–40 1303
 Triantaphyllou MV, Kouli K, Tsourou T, Koukousioura O, Pavlop- 1304
 oulos K, Dermitzakis MD (2010) Paleoenviromental changes 1305
 since 3000 B.C. in the coastal marsh of Vravron (Attiki, SE 1306
 Greece). *Quat Intern* 216:14–22 1307
 Vajsov I (1998) The typology of the anthropomorphic figurines from 1308
 Northeastern Bulgaria. In: Stefanovich M, Todorova H, Haupt- 1309
 mann H (eds) In the steps of James Harvey Gaul, vol 1. James 1310
 Harvey Gaul Foundation, Sofia, pp 107–141 1311
 Valamoti S (2004) Plants and people in Late Neolithic and Early 1312
 Bronze Age Northern Greece an archaeobotanical investigation. 1313
 BAR International Series 1258, Archaeopress, Oxford 1314
 Valsecchi V, Finsinger W, Tinner V, Ammann B (2008) Testing the 1315
 influence of climate, human impact and fire on the Holocene 1316
 population expansion of *Fagus sylvatica* in the southern Prealps 1317
 (Italy). *Holocene* 18:603–614 1318
 Velchev V (2002) Types of vegetation. In: Koprlev I (ed) Geography 1319
 of Bulgaria. ForCom Publishers, Sofia, pp 324–335 (in Bulgarian) 1320
 Wanner H, Beer J, Bütikofer J, Crowley TJ, Cubasch U, Flückiger J, 1321
 Goosse H, Grosjean M, Joos F, Kaplan JO, Küttel M, Müller SA, 1322
 Prentice IC, Solomina O, Stocker TF, Tarasov P, Wagner M, 1323
 Widmann M (2008) Mid- to Late Holocene climate change: an 1324
 overview. *Quat Sci Rev* 27:1,791–1,828 1325
 Weninger B, Alram-Stern E, Bauer E, Clare L, Danzeglocke U, Jöris 1326
 O, Kubatzki C, Rollefson G, Todorova H, Van Andel T (2006) 1327
 Climate forcing due to the 8200 cal yr B.P. event observed at 1328
 Early Neolithic sites in the eastern Mediterranean. *Quat Res* 1329
 66:401–420 1330
 Weninger B, Clare L, Rohling EJ, Bar-Yosef O, Böhner U, Budja M, 1331
 Bundschuh M, Feurdean A, Gebel H-G, Jöris O, Linstädter J, 1332
 Mayewski P, Mühlenbruch T, Reingruber A, Rollefson G, 1333
 Schyle D, Thissen L, Todorova H, Zielhofer C (2009) The 1334
 impact of rapid climate change on prehistoric societies during 1335
 the Holocene in the Eastern Mediterranean. *Documenta Prae-* 1336
historica 36:7–59 1337
 1338